

COLD SHORES IN WARMING TIMES – CURRENT STATE AND FUTURE CHALLENGES IN HIGH ARCTIC COASTAL GEOMORPHOLOGICAL STUDIES

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ABSTRACT. Many of the existing intellectual paradigms regarding the functioning of the polar coastal zone are now out-dated, based on descriptive geomorphology and a limited process-based understanding. Currently, among many components of Arctic landscape adjusting to global warming, the coastal zone is probably the most critical one both in terms of rapidity of environmental change as well as importance for human communities living in circumpolar regions. This issue was often raised during the 4th International Polar Year 2007–2008 and encouraged the scientific community to focus on the state of cold region coasts in more detail. In this paper I summarize the most recent developments in Arctic coastal geomorphology with a particular focus on the Svalbard Archipelago and draw attention to the research challenges awaiting further investigation. This paper highlights the need for a greater understanding of the controls on High Arctic coastal geoecosystems, especially given the potential for accelerated warming and sea-level rise in the coming decades and centuries. Many of presented views benefited from discussions with Professor Andrzej Kostrzewski – to whom this volume is dedicated.

KEYWORDS: cold region coastal geomorphology, global warming, sea-level change, High Arctic landscape dynamics, Svalbard.

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The coasts of neglect

Understanding the transition of Arctic regions, which in recent years have undergone a rapid environmental change in response to climate warming (Serreze *et al.* 2000; ACIA 2005; NorACIA 2011) is of crucial importance for the interpretation of Earth system reaction to former glaciations and climate shifts as well as more ac-

curate prediction of future scenarios of global climate change. The tremendous research efforts undertaken during the 4th International Polar Year 2007–2008 have revealed that the Arctic is already experiencing the strongest air and sea temperature rise on Earth. This process directly and indirectly changes the nature of the Arctic coastal zone. Of particular interest is the reaction of Arctic coastal geoecosystems to increased

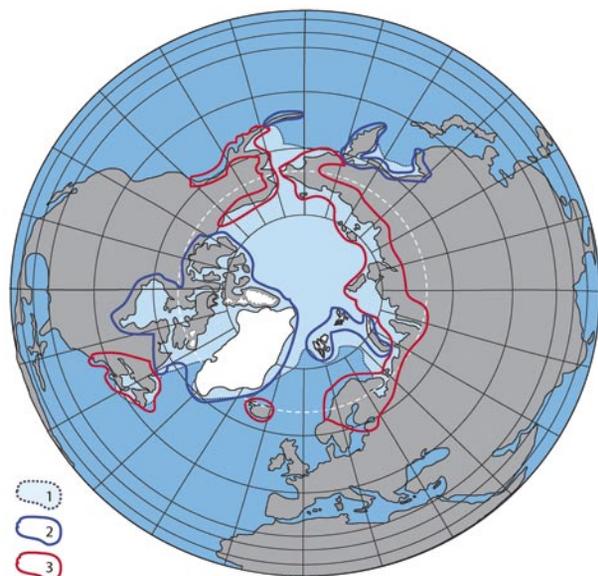


Fig.1. Simplified map of cold region coasts in the northern hemisphere (modified after Byrne & Dionne 2002).

1 - Extent of winter sea ice cover; 2 - Zone of still relatively poorly studied cold region coasts where the mechanisms controlling coastal evolution remain unknown; 3 - Zone of relatively well-studied cold region coasts with running coastal erosion monitoring programmes, and number of high quality papers published in a last decade.

sediment and nutrient supply from land as well as increasing open-water conditions leading to intensified shoreline erosion and sediment transport. The recently published report 'The State of Arctic Coast 2010 Report' (Forbes *et al.* 2011) suggests that circumpolar coastal zone (Fig.1) is the key interface in the entire Arctic characterized by 'most rapid and severe environmental changes which have serious implications for communities living on coastal resources'. However, many of the existing opinions regarding the functioning of polar coasts based on limited observations are invalid and reduces our ability to understand current mechanisms controlling the present state of this fragile zone and do not deliver any sort of prediction of its future evolution. Not only is the number of academic papers on high latitude coastal environments lower than from temperate and tropical regions, but also their qualitative nature seems to be insufficient to allow numerical modelling and more sophisticated data analyses. In contrast to lower latitude coasts little is known regarding the potential impact of climate and sea-level change on high latitude coastal margins. This relative lack of research advance has persisted for several decades. This is highlighted by the remark from Trenhaile (1983) that 'there

is a lack of even basic agreement on the efficiency of coastal processes in the high latitudes' then almost twenty years later Byrne & Dionne (2002) reviewing the state of high latitude coastal geomorphology still had to admit that coastlines, accounting for at least 30% of the world coasts, have been neglected and studies on their nature are very limited. Most recently Lantuit *et al.* (2011) indicated that only 1% of Arctic coastlines have been investigated in sufficient detail to allow quantitative description of processes operating on them. One of the key messages found in all major papers reviewing the developments in cold region coastal studies is the need for long-term coastal change monitoring and selection of core geo-indicators which could be analyzed in a unified way under the umbrella of a circumpolar observation network (e.g. John & Sugden 1975, Forbes & Taylor 1994, Trenhaile 1997, Byrne & Dionne 2002, Urdea 2007, Forbes *et al.* 2011). Though knowledge of the response of the coastal zone to climate change is fundamental to understanding polar landscape adjustment to global warming, we still cannot precisely measure the rate and scale of changes for vast areas of the Arctic. At the heart of the problem lies a mismatch between out-dated concepts based on short field observations and the lack of the studies linking Arctic coastal evolution with transformation of Arctic landscape occurring during climate shifts on millennial (glacials/interglacial), decadal (e.g. changes in atmospheric circulation patterns such as NAO/AO) and seasonal (i.e. the duration of snow cover, high/low precipitation, high/low air temperatures, strong/weak winds etc.) timescales.

'Known knowns, known unknowns and unknown unknowns'

Cold region coasts are typically defined as 'those areas where frost and ice processes are active during a period of the year which is sufficient to have a significant, if not permanent, impact on the near terrestrial, coastal and marine environments' (Byrne & Dionne 2002). Several authors (e.g. Forbes & Taylor 1994, Trenhaile 1997) noticed that many morphological aspects (cliffs, beaches, barriers, dunes, spits, embayments, lagoons) and



Fig. 2. High Arctic coastal environments in northern Billefjorden:

1 - sandstone boulder-strewn beach in northern Adolfbukta; 2 - plunging metamorphic rock cliffs in northern Adolfbukta; 3 - weathered rock cliffs and mixed sand-gravel barrier in eastern Petuniabukta; 4 - gravel barrier in eastern Petuniabukta; 5 - tidal flat (right), gravel spit (middle) and lagoon (left) in northern Petuniabukta; 6 - spit platform developing at the mouth of Ebbaelva (background), at the front uplifted beach sequences in NE Petuniabukta; 7 - tidal creek in central part of tidal flat (northern Petuniabukta); 8 - High Arctic salt marsh in NE Petuniabukta.

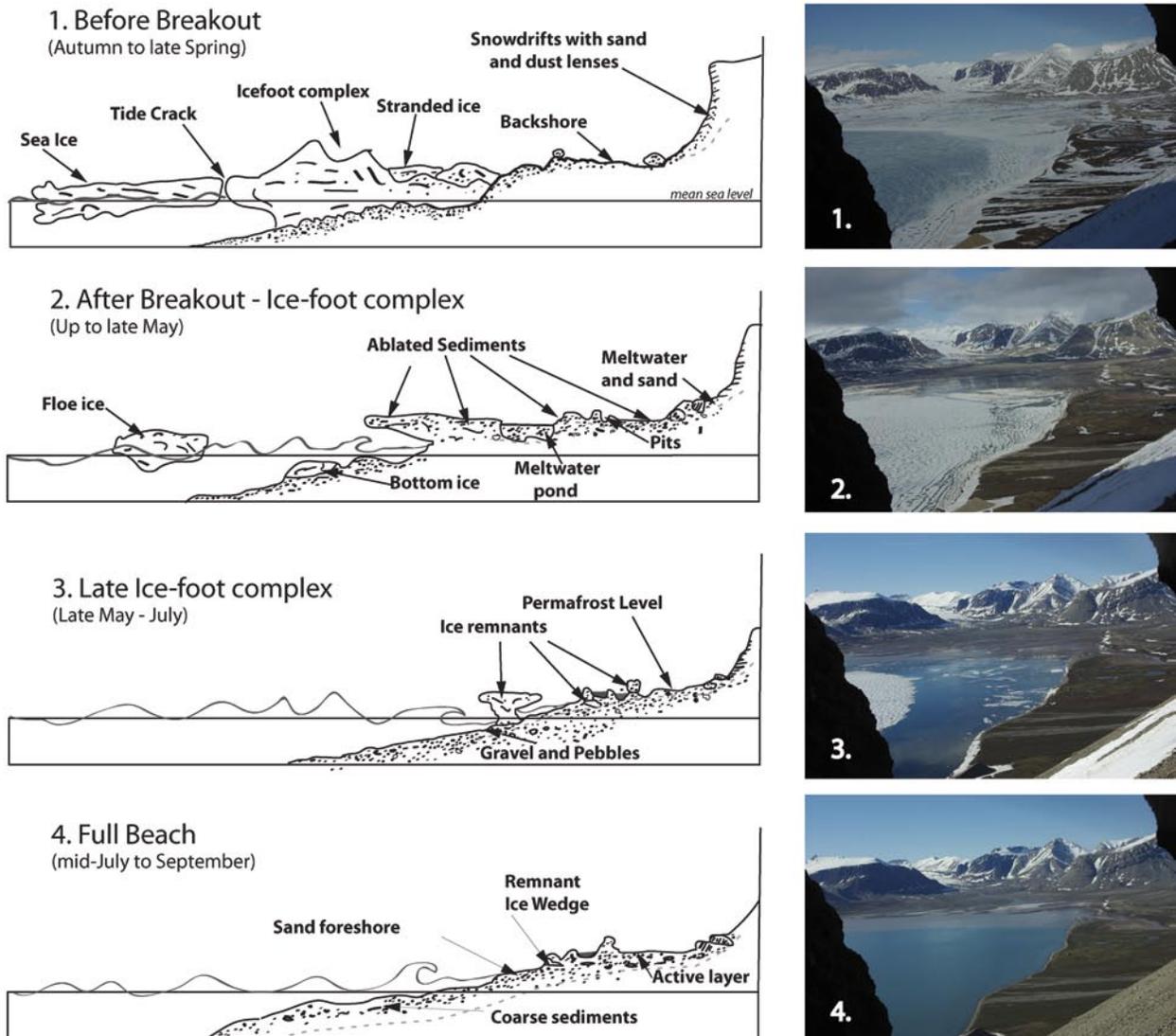


Fig. 3. Seasonal changes in polar beach profiles (sketches after Hansom and Kirk 1989) as they adjust to open water conditions. On the right examples of different stages of beach evolution from Petuniabukta (images taken by time-lapse camera installed on Wordiekammen massif in years 2009–2010)

processes (tides, currents, waves) operating on coasts in cold environments (Fig. 2) are similar to those from lower latitudes. What distinguishes them, however, are the effects of permafrost, ground ice, frost action, sea ice, snow cover, isostasy, glacial history and other zonal factors on coastal morphodynamics. It is important to notice that the scientific description of cold region coastal environments is characterized by many discrepancies. For instance, tremendous effort of Canadian researchers in last 40 years brought a major advance in our understanding of processes controlling Arctic tidal flats and development of the term *glacielto* to describe ice-influenced landforms, sediments, processes in marine, coastal,

estuarine, lacustrine and fluvial environments (e.g. Dionne 1968, 1969, 1988, 1989; Drake & McCann 1982; McCann *et al.* 1981; McCann & Dale 1986; Dale *et al.* 2002). Also Arctic beach environments (Fig. 3) have been solicitously described in many sites (e.g. Rex 1964; Hume & Schalk 1964a, 1964b, 1967, 1976; Zenkovitch 1967; King & Buckley 1968; McCann & Owens 1969; Green 1970; Owens & McCann 1970; McCann & Taylor 1975; Rosen 1978; Taylor 1978; Reinson & Rosen 1982; Sempels 1987; Reimnitz *et al.* 1990; Barnes *et al.* 1993, Campeau & Héquette 1995) whereas to date there have been relatively few studies investigating Arctic rocky coasts (e.g. Jahn 1961; Nielsen 1979; Dionne & Brodeur 1988, Fournier

& Allard 1992, Ødegård & Sollid 1993; Ødegård *et al.* 1995; Lundberg & Lauritzen 2002; Wagenseen *et al.* 2007; Strzelecki 2011). This gap is even more profound if we take into consideration that up to 35% of Arctic coastlines are rock-dominated (Lantuit *et al.* 2011). The classic example showing the research stagnation in revealing the complexity of high latitude rocky coastal zone is the longstanding controversy regarding the origin of strandflats. This has been a regular topic for discussion in the geomorphological literature for almost a century (Nansen 1922; Guilcher *et al.* 1986; Holtedahl 1998) and a debate on the efficiency of periglacial shoreline evolution, with enthusiasts of slow (Zenkovich 1967) and relatively rapid shore evolution (Jahn, 1961).

Furthermore, the development in Arctic coastal geomorphology is marked by significant regional contrast. The last decade has seen major developments in Arctic coastal research due to research from the Arctic Coastal Dynamics (ACD) Group (Rachold *et al.* 2005; Overduin *et al.* 2007; Lantuit & Pollard 2008, Lantuit *et al.* 2011). Another important step was the reopening of Russian works to the wider scientific community, especially in the field of thermoabrasion and coasts formed in the Yedoma formation (e.g. Aré 1988; Nikiforov *et al.* 2005; Leontiev 2006; Aré *et al.* 2008; Lukyanova *et al.* 2008; Streletskaia *et al.* 2009). The major focus in these initiatives has been the understanding and modelling of ice-rich permafrost along Alaskan and Siberian coastlines. The study by Lantuit *et al.* (2011) indicated that in comparison to Beaufort and Laptev Sea coastlines, the coasts of High Arctic archipelagos such as Svalbard, Franz Josef Land, the Canadian Arctic Archipelago (CAA) and Greenland, whose melting ice masses contribute the most to present-day sea level rise, are still poorly recognized. Nevertheless, thanks to research projects carried out by ACD (Fig. 4) we can consider the beginning of the 21st century as a *golden decade* of Arctic coastal studies. Their main findings and predictions from the perspective of coastal geomorphology are:

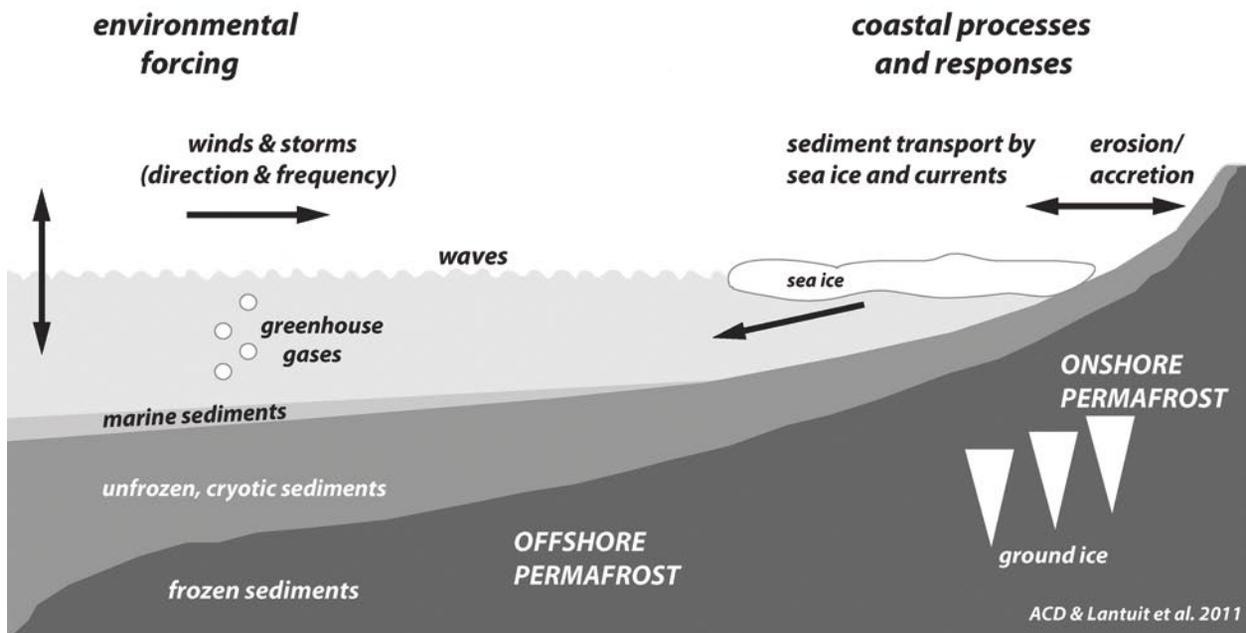
- the expected strong link between Arctic coastal evolution in coming decades and climate warming;
- the duration, intensity and number of storms entering the Arctic has already increased and

the wave influence on the Arctic coasts will continue to intensify with implications for accelerated erosion;

- the changing Arctic sea-ice regime and the timing of ice break-up and freeze-up may result in further reduction of coastal protection from waves (in last decade sea-ice extent reached record minima in extent and perennial ice has been reduced by almost 50% in extent);
- the increased sea-surface temperatures, longer ice-free periods and intensified thawing of subsea, coastal and inland permafrost leads to rapid erosion and retreat of Arctic coasts as it was documented in recent years in Alaska and Siberia. ACD calculations of mean rates of coastal retreat in the last decades are typically in the 1–2 m/year range, but vary up to 10–30 m/year along the Beaufort Sea, the East Siberian Sea, and the Laptev Sea. ACD proved that previous reports underestimated the sediment and carbon supply from Arctic coasts to the ocean;
- the existence of a strong link between sea-level rise in the Arctic coastal zone and warming of the coastal waters as well as increased water freshening;
- the rise in sea level is expected to enhance coastal erosion and affect sediment transport in coastal areas (various IPCC models show higher sea-level rise in the Arctic than mean-global rise);
- glacial legacy (glacioisostatic uplift/subsidence) is a dominant factor controlling the relative sea-level change and is characterized by strong regional contrasts e.g. Canadian Arctic Archipelago with emerging central part and submerging margins;
- climate warming leads to increased Arctic river discharge and sediment supply to the coastal zone. However, in some locations increased terrestrial sediment delivery is not sufficient to balance the negative changes caused by sea-level rise and storm flooding, what leads to submergence and erosion of Arctic deltas, e.g. the second largest Arctic delta, the Mackenzie Delta, is retreating dozens of meters per year.

Currently, the leaders of ACD are planning the organisation of the Arctic Circumpolar Coastal Observatory Network, which should cover more

A) PROCESSES ACTIVE ON ARCTIC PERMAFROST COASTS (e.g. SIBERIA)



B) PROCESSES ACTIVE ON HIGH ARCTIC PARAGLACIAL COASTS (e.g. SVALBARD)

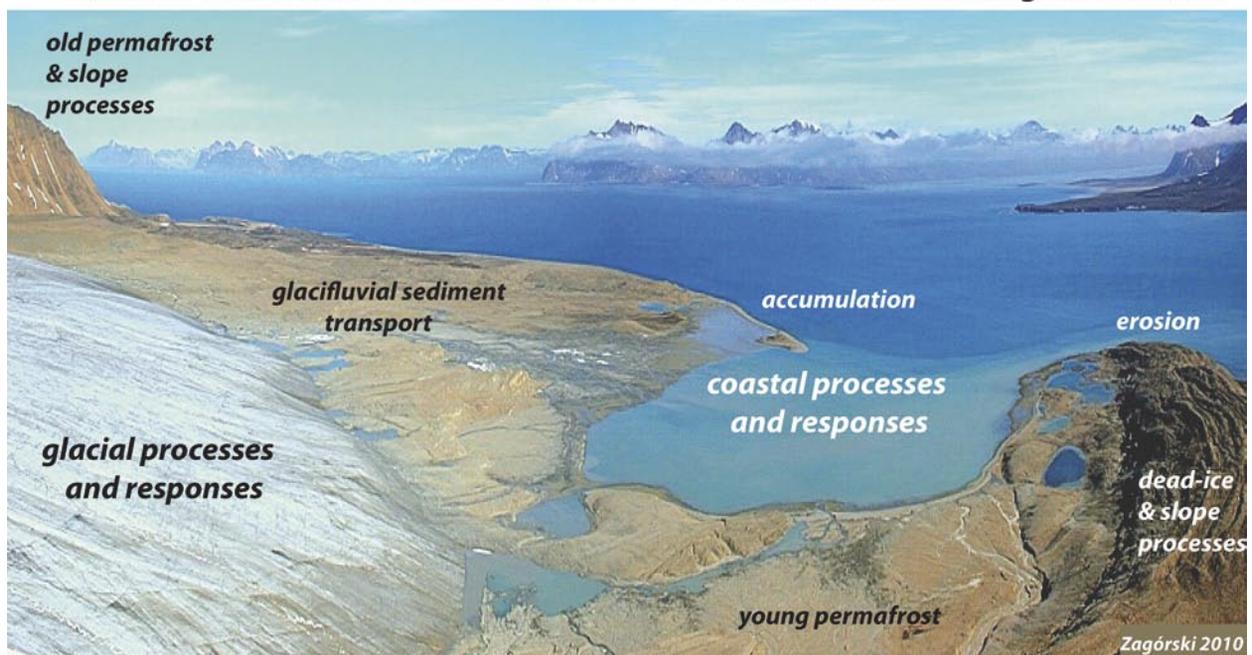


Fig. 4. Functioning of Arctic coastal zone:

A - model developed by Arctic Coastal Dynamics Group for ice-rich permafrost coasts characteristic for Siberian and Alaskan coasts (after Lantuit *et al.* 2011); B - the complexity of High Arctic coastal adjustment to glacier retreat - no model exists (Photo by Piotr Zagórski, 2010, illustrating paraglacial coastal system in Bellsund, Spitsbergen).

sites in the High Arctic and diametrically change the regional disproportion. It is then important to ask what kind of research challenges await the scientific community along the High Arctic coasts.

High Arctic enigma

To date coastal research in High Arctic settings has been strongly linked to the successful application of the paraglaciation concept (*sensu*

Slaymaker, 2011) in cold region geomorphological studies. During the 1990s French geomorphologists dominated research on High Arctic paraglacial coastal systems. Their research focused on the response of the coastal zone to changes in glacial systems. Héquette & Ruz (1990) and Héquette (1992) suggested that although High Arctic coasts are thought to be low energy in a certain conditions they can transform very rapidly. Their study on coastal margins of glaciofluvial outwash plains in northwest Spitsbergen (Brøgger Peninsula) documented a whole mosaic of abrupt changes in coastal morphology during the short Arctic summer with coexisting episodes of erosion and deposition, formation of new and cannibalization of old barriers, deposition of micro-deltas, landward and seaward migration of coastlines depended on rate of glaciofluvial activity in the surrounding and ability of storm waves to overwash barrier crests.

Mercier & Laffly (2005) linked the periods of intensified sediment supply, associated with the post-Little Ice Age (LIA) retreat of Pedersenbreen, Austre, Midre and Vestre Lovénbreen with coastal progradation. During the last 30 years, the mean annual coastal progradation along studied section of Kongsfjorden was 3 m. It has also been noted that in coastal sections where glaciofluvial sediment delivery was reduced shoreline recession was observed. This study demonstrated the high sensitivity of Svalbard coasts to mechanisms controlling the supply of sediments from deglaciated catchments. In this respect their work confirmed earlier observations from Atlantic Canada, made by Forbes and Taylor (1987) who recorded on-shore migration of gravel barriers during periods of sediment shortage and spit extension after an increase in sediment input. In more sheltered area of Spitsbergen, centrally located – Petuniabukta, characterized by limited fetch and prolonged sea-ice conditions, Strzelecki (2011) reported similar rates of coastal adjustment. This research suggests that High Arctic coasts fed by paraglacially reworked and transported sediments deserve a new model, which will properly reflect the interplay between sediment supply, sea-level change and glacial landscape transition. This is important because many sections of the Svalbard coast are clearly still responding to a combination of ongoing paraglacial and coastal processes.

In recent decades paraglacial processes have become one of the most effective geomorphological processes on Svalbard, reducing the impact of glacial processes to a secondary role in landscape transformation (Laffly & Mercier 2002, Rachlewicz 2010). Also the previous natural shoreline protection in the form of sea-ice, ice-foot complex and snow cover now disappear earlier and form later in the season allowing warmer and stormier sea conditions to access unconsolidated shorelines.

Based on several environmental monitoring case studies on Strokdammene Plain (W Spitsbergen) Åkerman (2008) proposed a geoecological model of Svalbard coasts under global warming conditions (Fig. 5). Åkerman predicts that prolonged duration of the open waters will be critical to the future impacts of coastal and near-coastal processes and environments in sites such as Svalbard. With frequent storms abrading unprotected shores additionally destabilized by thawing permafrost, flooding of coastal lowlands, high and wide barrier systems blocking stream outlets, Svalbard coastal geocosystem will change rather abruptly.

On the other hand, Ziaja *et al.* (2009) already documented a whole cycle of coastal landscape transition related to the post-LIA deglaciation of Humbergbukta in SE Sorkapp which highlighted the role of changes in the High Arctic coastal zone as a simulator of polar life expansion. Retreat of Hambergbreen and other surrounding glaciers not only led to formation of new coastlines which quickly become transformed by coastal processes, but also exposed new areas for plant succession and animal colonization.

Another example of Svalbard coastal zone adjustment to climate warming and change in sediment delivery pathways was a study on the decadal-scale shoreline dynamics of Calypsotranda carried by Zagórski (2011). His investigations, combining remote sensing and field-based observations indicated the dominance of coastal erosion over accumulation with maximum coastal retreat up to 100 m, regardless of the retreat of surrounding glaciers and increased sediment supply to the fjord system. In addition, based on a study of post-LIA geosuccession in Recherfjorden region Zagórski *et al.* (in preparation) developed a concept of direct and indirect glacial in-

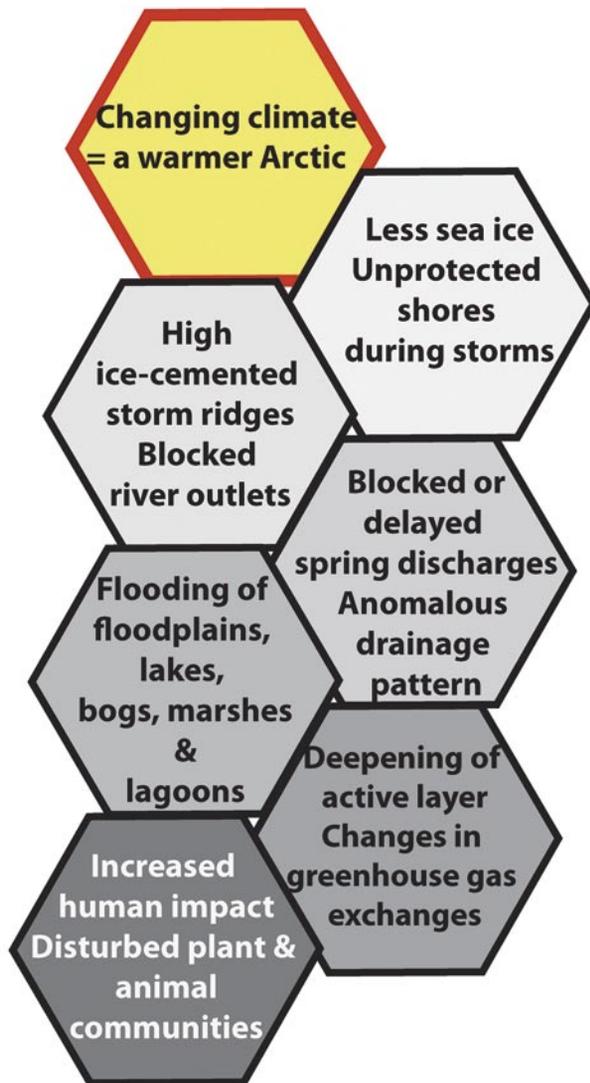


Fig. 5. Scenario of geocological changes of Svalbard coasts under global warming conditions (modified after Åkerman 2008).

fluence on High Arctic coastal evolution which differentiates coasts formed after retreat of tide-water glaciers from coasts formed due to intensification of sediment delivery from valley-glacier systems. The direct influence of glacier systems on the coastal zone was previously documented by Nielsen (1992) in West Greenland. The retreat of the tide-water glacier Equip Sermia after the 1920s advance led to the transformation of a lateral moraine into a barrier spit and formation of a lagoon system in just 70 years. Nielsen argued that such a rapid landform transformation has serious implications for Holocene landscape interpretation, especially in case of landform and sediment assemblages related to tide-water glacier systems.

He also postulated a role of high waves induced by ice-berg roll from calving in over-washing and reshaping High Arctic coastal barriers, leading to the formation of boulder barricades and beaches. It has to be stated that the role of landslide (paraglacial) or ice-berg induced tsunami waves on the Arctic coastal morphology has never been studied in detail. Later on, Nielsen (1994) expanded his concept of glacier/coastal interaction during the study of Arctic delta in Sermilik Fjord. This is a key work documenting the sensitivity of the high latitude coastal zone to post-LIA landscape transition from glacier into paraglacial dominated geocosystem and the rapidity with which such change can occur. During one century the delta system experienced all stages of evolution from rapid formation after retreat of Mitdluagkat glacier; progradation in response to intensified sediment supply; development of spits, barriers, tidal island and tombolos; finally to a degradation by waves and currents under rising sea-level conditions after reduction of delivery from glacier catchment. More recently Hansen (2004) illustrated the complexity of a High Arctic delta system response (Hall Bredning-Scoresby Sund, East Greenland) to change in topography, sediment supply, rate of deglaciation and shifts in climate on the Holocene timescale. Hansen's model can be adapted to hundreds of small and sheltered Arctic fjords deglaciated during the Holocene; it also proves that High Arctic coastal environments may serve as a useful tool in reconstructing environmental change.

Stories to be rewritten

One of the classic examples of the application of sedimentological survey of coastal landforms in palaeogeographical reconstruction is a study by Lønne & Nemeč (2004) on High Arctic coastal fan delta in Advenfjorden (Spitsbergen). By detailed sedimentological analysis of fan depositional structures they demonstrated the importance of fan deltas as a proxy record of deglaciation history and environmental change in a High Arctic terrain, especially important in regions characterized by low preservation of glacial sediments and landforms. The small fan system developing at the foot of Hiorthfjellet was a very precise re-

order of the interplay between terrestrial and marine processes adjusting to shifts in sediment delivery from retreating glacier as well as fluctuations of a sea-level. This study, apart from documenting in an unprecedented detail the sedimentology of a polar fan system fed by glacier streams, snow melt and degrading permafrost, also provided the first well-dated onshore proof for a Mid-Holocene transgression in central Spitsbergen. Surprisingly, even after this important research, the significance of cold region coastal fan systems still remains unrecognized.

Another important story can be added to reconstruction of polar landscapes by studying the sedimentology of gravel-dominated storm ridges preserved in many uplifted marine sequences across High Arctic region. In Andréeland (NW Spitsbergen) Brückner *et al.* (2002) and Brückner & Schellmann (2003) postulated that uplifted beach ridges may help us decipher the glacio-isostatic and palaeoceanographic changes. They based their hypothesis on the assumption that conditions favourable for formation of beach ridges reflects existence of ice-free conditions associated with retreat of glacier systems and open-water summer season with wave and current activity. In the case of Woodfjorden and Wijdefjorden, formation of hundreds of Holocene beach ridges was possible only due to climate warming which lead to retreat of the glacier, opening of the fjord and entrance of warm waters brought by the West Spitsbergen Current securing removal of sea ice and open water conditions in summer months. More recently St-Hilaire-Gravel *et al.* (2010) studied the topography and sedimentology of uplifted beach ridges on Lowther Island (CAA) in order to test their ability to record past wave intensity and sea-ice conditions. However, reflecting only local conditions they suggested that better-developed beach ridges can be associated with periods of prolonged wave activity (increased duration of ice free periods). One future direction of research might be to compare the onshore record (beach ridges, coastal fan deltas) with the offshore record from ocean cores collected across the High Arctic to shed new light on the history of sea ice and ocean currents.

Action instead of conclusions

I strongly believe that the described gap in understanding of High Arctic coastal adjustment to global warming conditions represents a great opportunity for the national polar research community. Longstanding history of Polish investigations on Svalbard with environmental monitoring programmes running for decades as well as a network of research stations spread across Spitsbergen constitute a solid basis for an interdisciplinary research project focusing on the functioning of the coastal zone in the heart of the European Arctic. Due to its location at the boundary between oceanic and atmospheric fronts, Svalbard is one of the key areas to study the Arctic sensitivity to climate change. The Svalbard region has, therefore, been an area of many scientific breakthroughs including: palaeo-climate reconstructions (e.g. Isaksson *et al.* 2005); the extent of the last glaciation (e.g. Mangerud *et al.* 1998; Svendsen *et al.* 2004); Holocene sea level changes (e.g. Forman *et al.* 2004); modern and relict glacial systems (e.g. Boulton 1972; Lønne & Lyså 2005; Nuth *et al.* 2010); periglacial permafrost processes and mechanisms (e.g. Humlum *et al.*, 2003); ocean water interaction and mixing (e.g. Ślubowska-Woldengen *et al.* 2007, Majewski *et al.* 2009). Despite the wealth of research which has been undertaken in Svalbard, there remains a lack of investigation into coastal processes and environments. Previous coastal studies (Mercier & Laffly 2005, Ziaja *et al.* 2009, Strzelecki 2011, Zagórski 2011) document dramatic changes in sediment flux and coastal response under intervals characterised by a warming climate, retreating local ice masses, a shortened winter sea-ice season and melting permafrost. These terrestrial processes are interacting with glacio-isostatic land emergence and on-going global sea-level rise. The pristine coasts of Svalbard provide a superb opportunity to quantify how High Arctic coasts are responding to rapid climate warming. Nevertheless several crucial research questions are still to be answered including: *What processes control existing Svalbard coasts?; How sensitive is the Svalbard coastal zone to recent increased human impact?;*

What was the response of the Svalbard coastal zone to intensified glacial recession and sediment supply after the Little Ice Age?; Do coastal processes intensify the rate of rock weathering on Svalbard climates?; How did the Svalbard coast and continental shelf respond to climate change, glacier fluctuations and sea ice changes throughout the Holocene?; What is the precision of the morphological and sedimentological records of environmental transformations, climate change and natural disasters along the Svalbard coast?

These research premises and challenges led to the formation of a SVALCOAST research team linking geoscientists from leading national polar research institutions and international scientific partners which in coming years will aim to examine the response of the coastal zone of the Svalbard Archipelago to current climatic warming and multidirectional anthropopression and explore the changes experienced by the coastal landscape during various stages of deglaciation. As well as this, predictions will be made as to the coast's future evolution under scenarios of continuing increases in air temperature and sea-level rise (Strzelecki & Zagórski 2011). Given the diversity of research questions and rapidity of coastal adjustment it seems to be a fascinating challenge!

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Fig. 6. Prof. Andrzej Kostrzewski and Matt Strzelecki during geomorphological mapping of western coast of Petuniabukta, Spitsbergen August 2007.

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